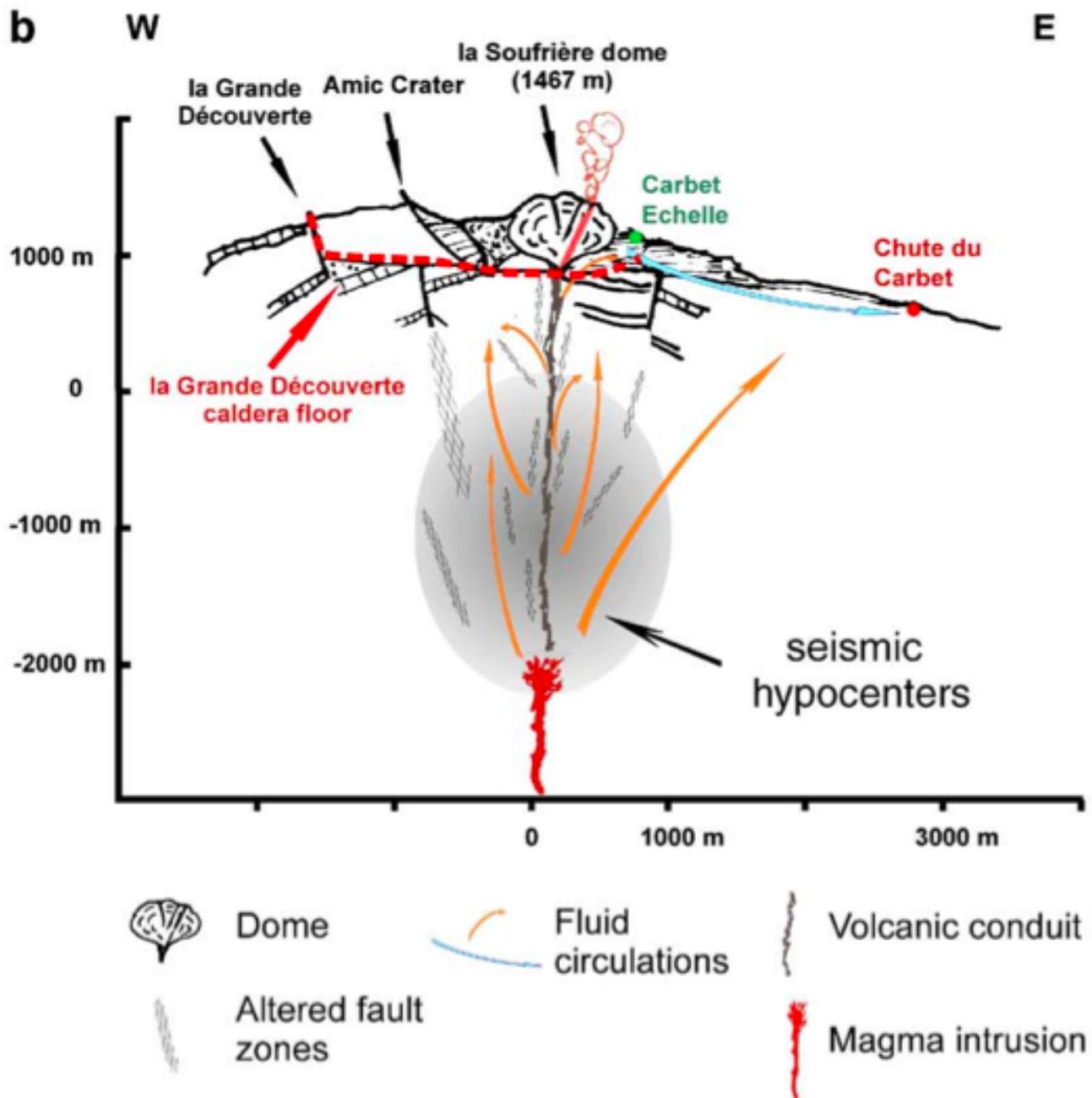
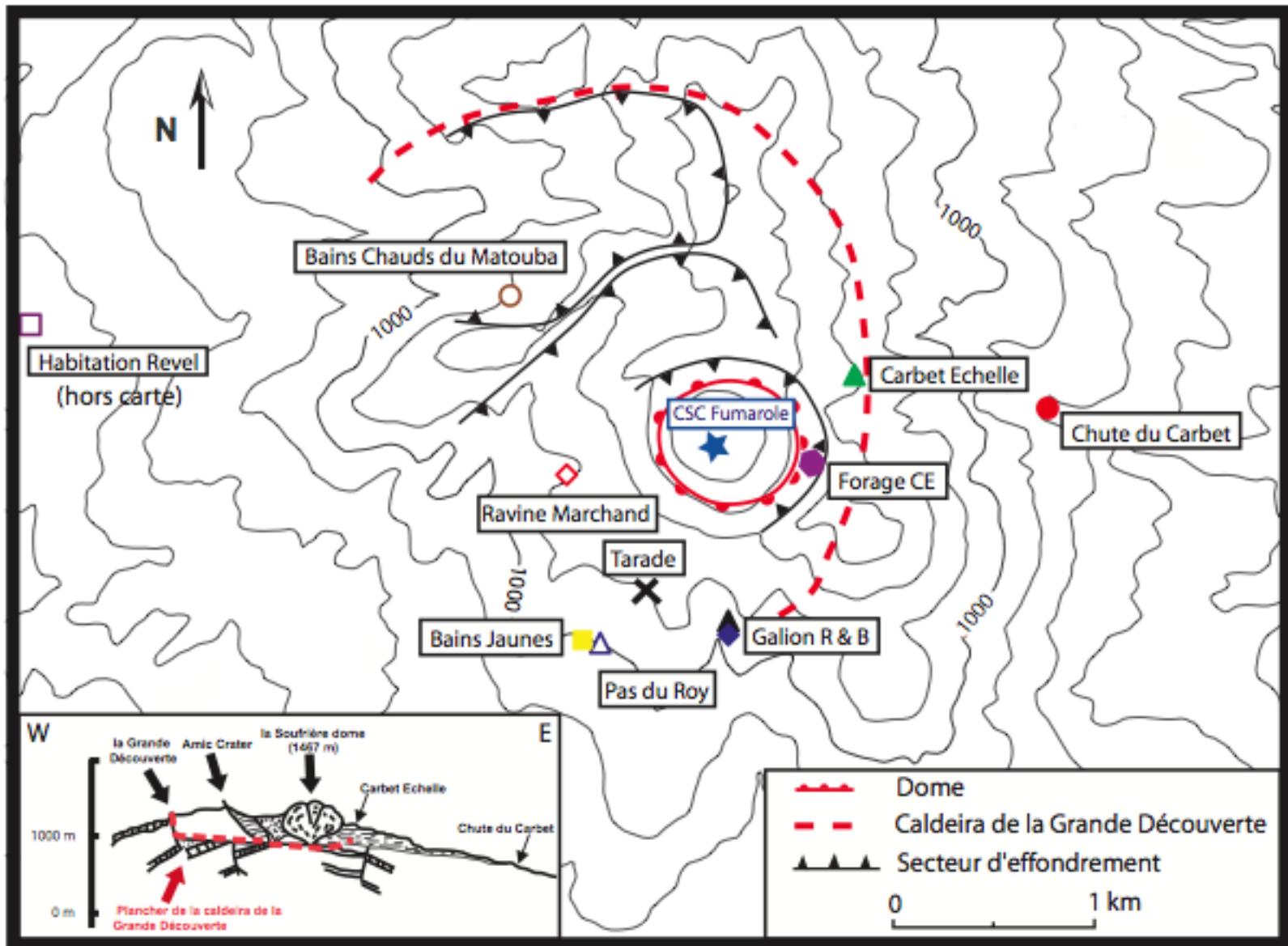


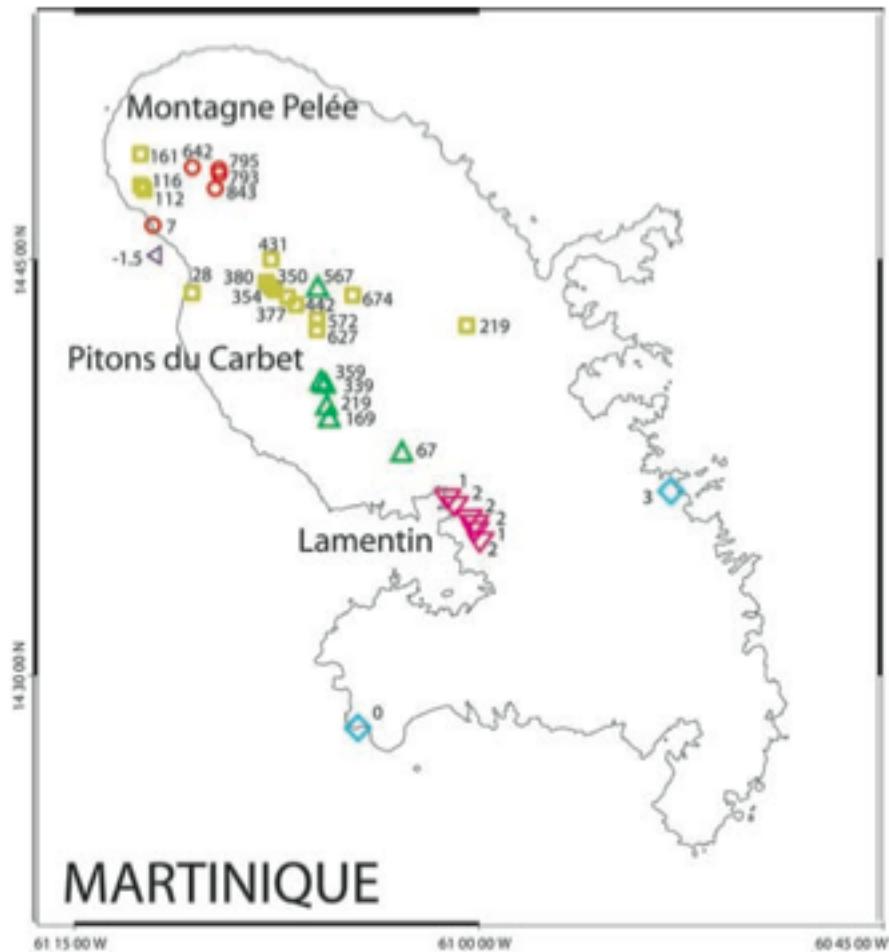


Le système hydrothermal de la
Soufrière de Guadeloupe et de la
Montagne Pelée de Martinique

- Gilbert Hammouya
- Olivier Crispi
- Michel Feuillard
- Jean-Christophe Komorowski
- Jean-Louis Cheminée
- Céline Dessert
- François Beauducel
- Thierry Kitou
- Christian Lambert
- Christian Antenor-Habacz
- Arnaud Lemarchand
- Michel Semet
- Gil Michard
- Cyril Aubaud
- Nathalie Jendrzejewski
- Long Li
- Lorraine Ruzié
- Manuel Moreira
- Magali Bonifacie
- Carine Chaduteau
- Jean Paul Toutain
- Benoit Villemant
- Agnès Michel
- Georges Boudon







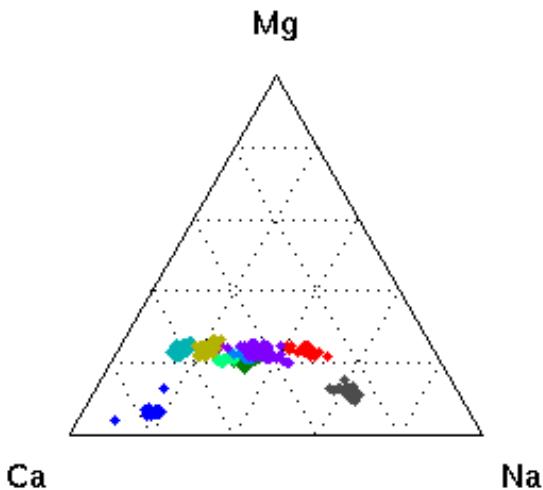
Synthèse Réseau Analyse Sources Thermales (10 ans)

08-Mar-2012 11:37:12 -4 - État 089 % - Acquisition 101 % - M - 1 mois

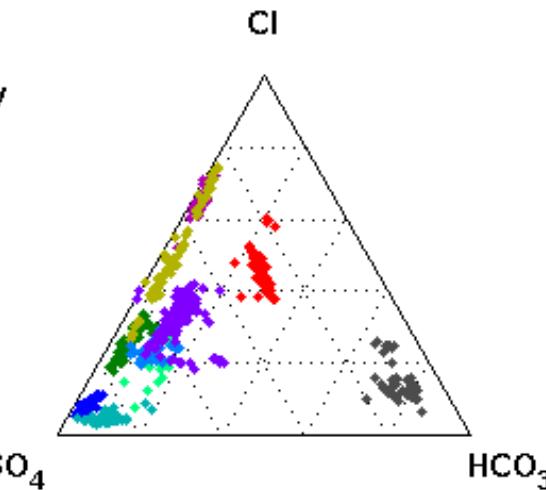
WEBOBS © 2012, OVSG-IPGP



Mg



Cl



Guadeloupe

Synthèse Réseau Sources Thermales (10 ans)

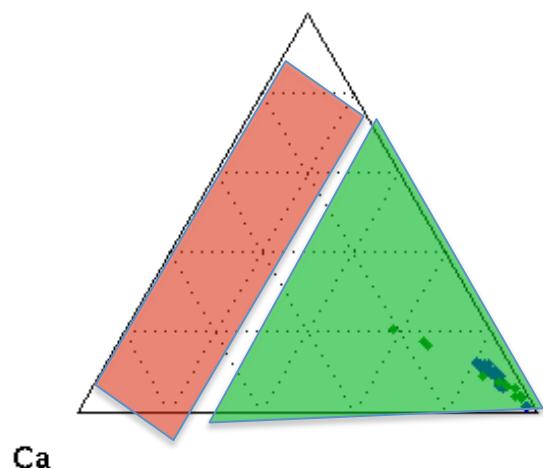
23-Dec-2010 02:43:56 +0 - État 000 % - Acquisition 046 % - M - 1 mois

WEBOBS © 2010, OVSG-IPGP



Mg

Cl

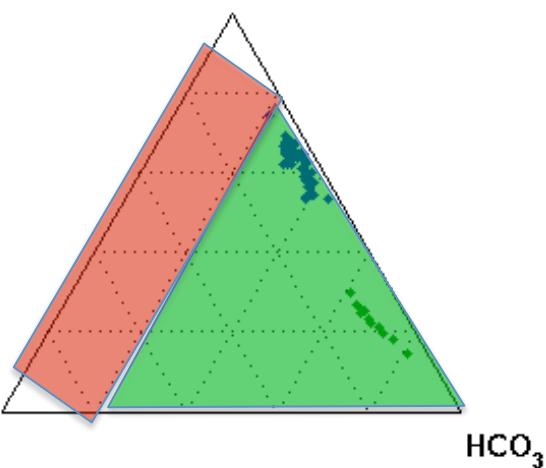


Ca

Na

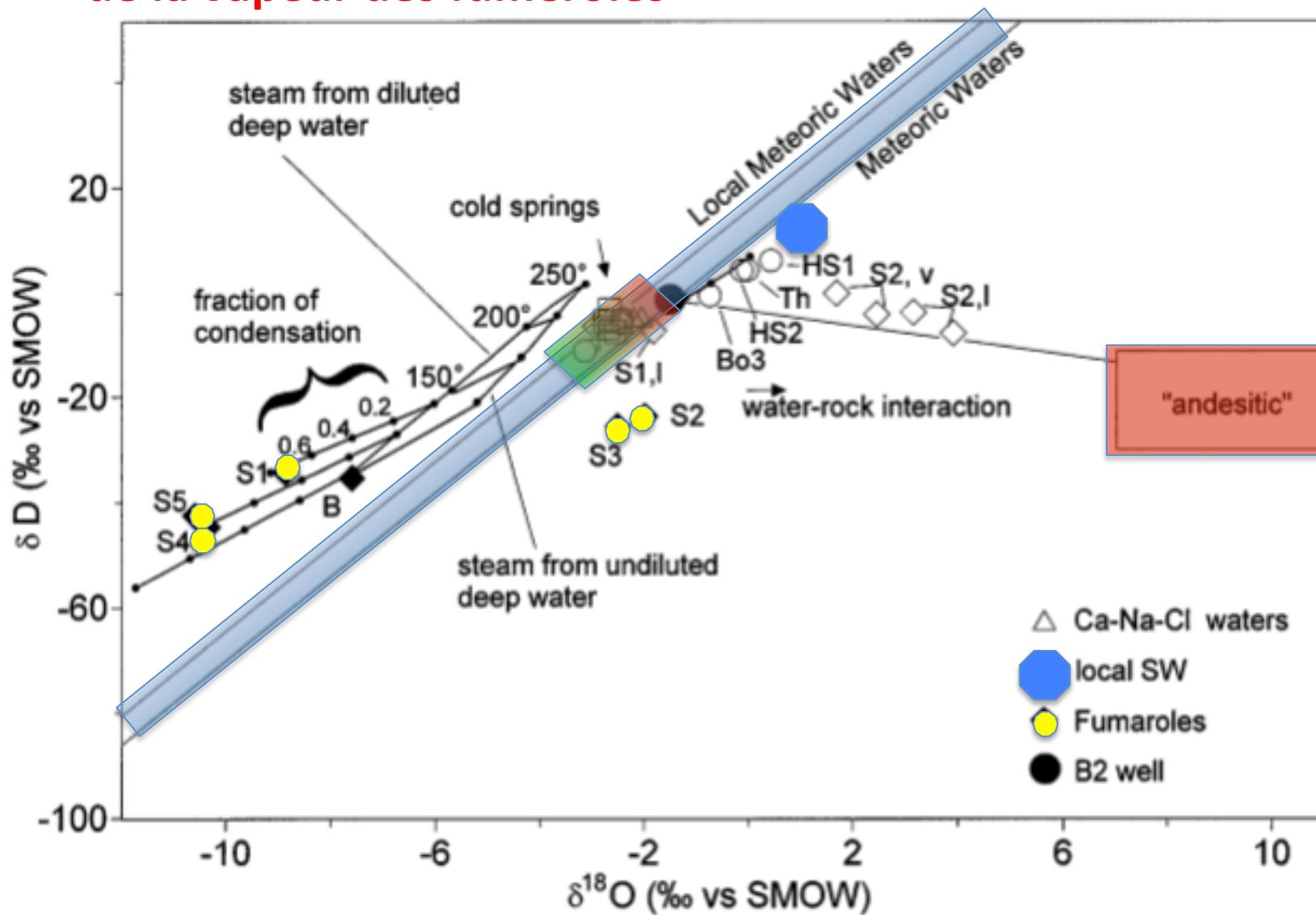
SO₄

Martinique

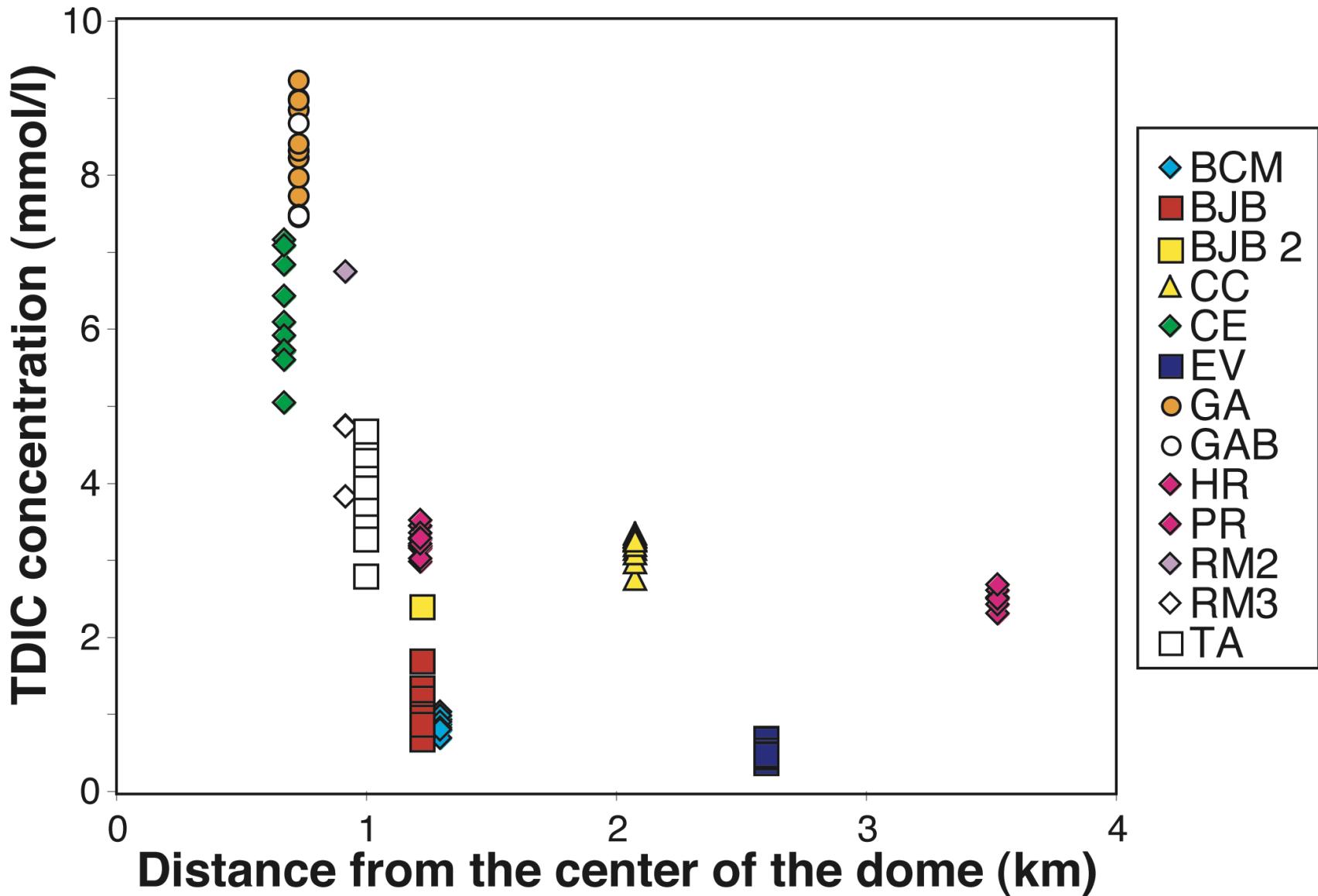


HCO₃

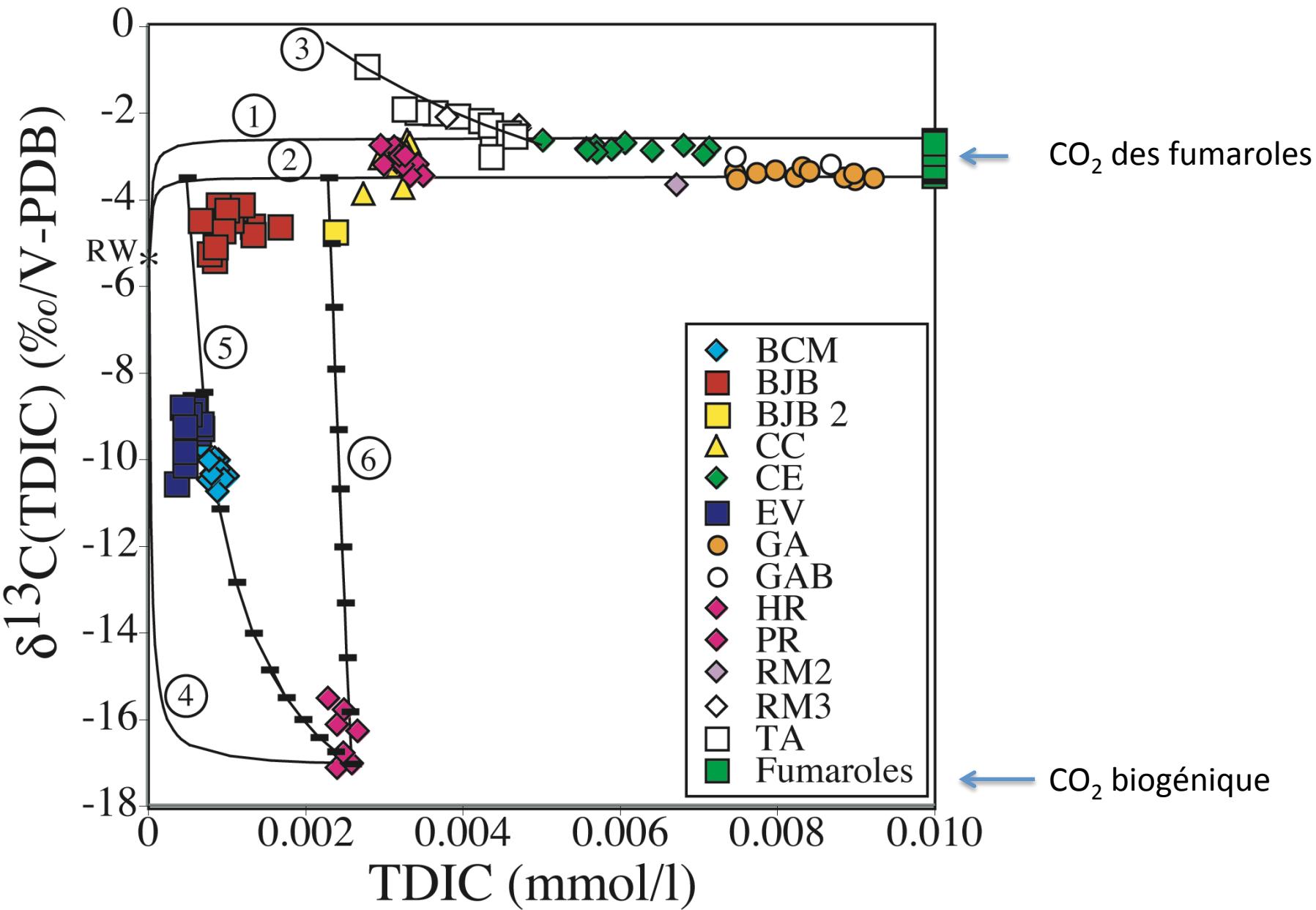
Composition isotopique de l'eau des sources et de la vapeur des fumeroles



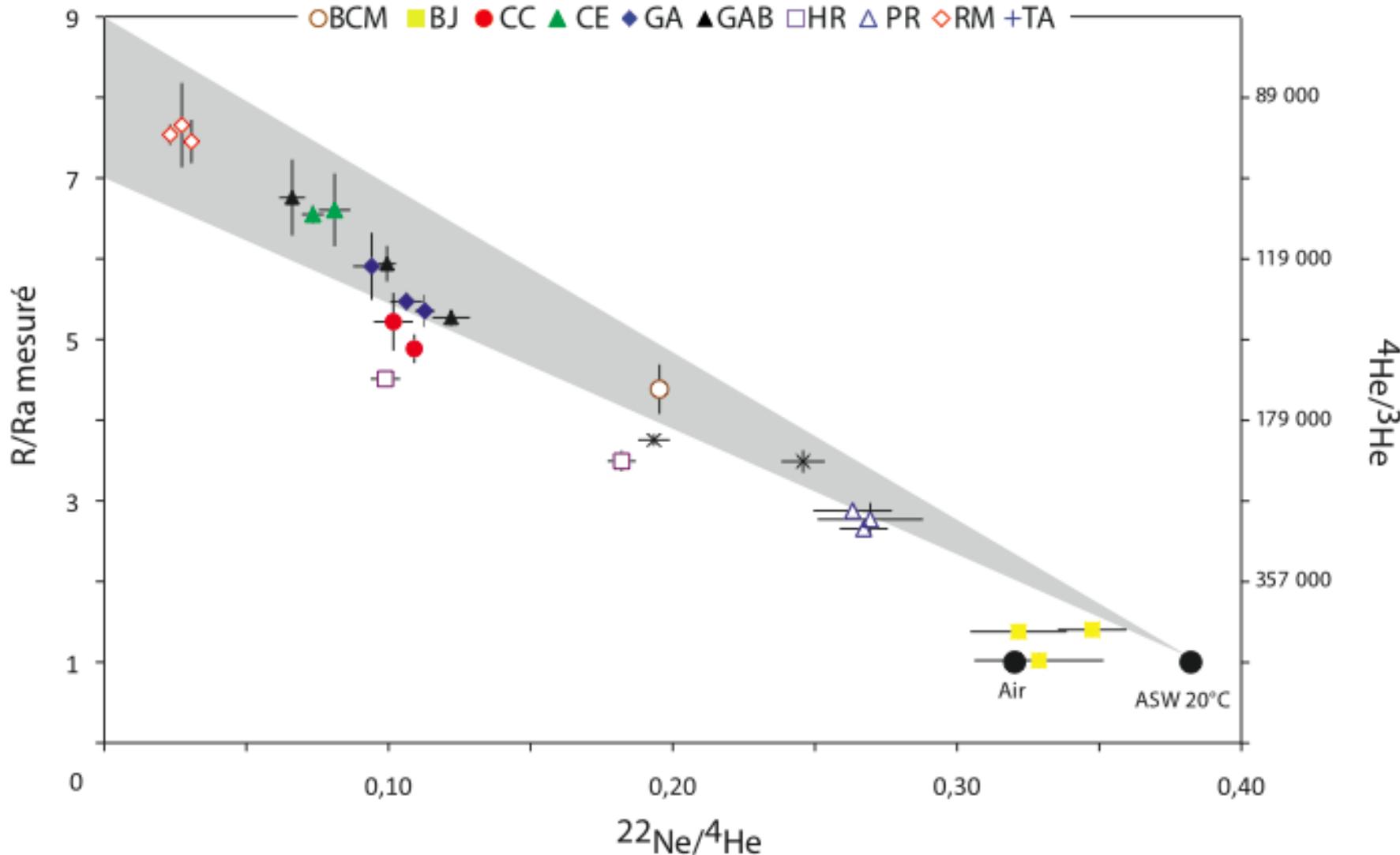
CO₂ dissous dans les eaux de sources

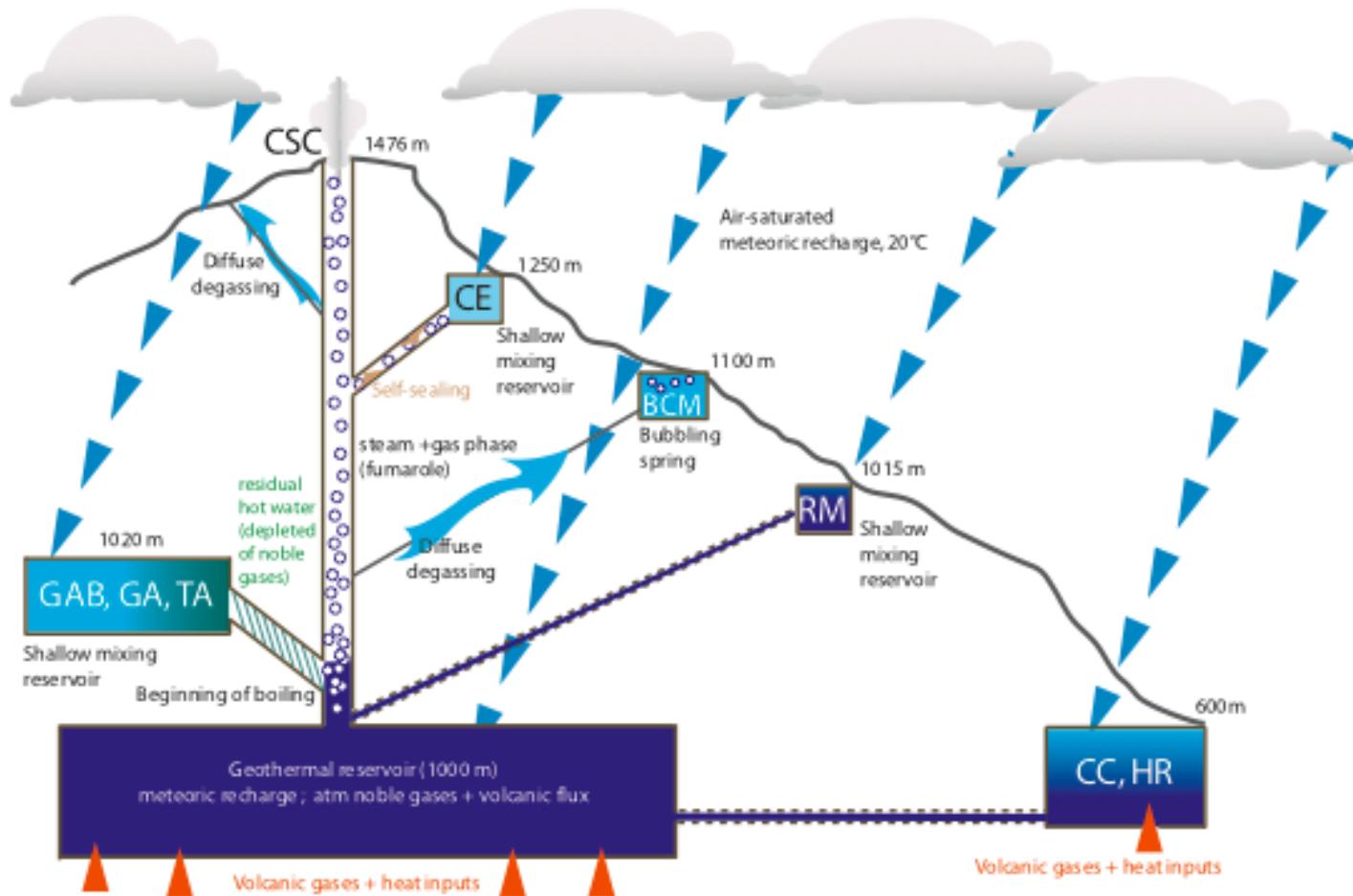


D'où vient le CO₂ dissous dans l'eau des sources



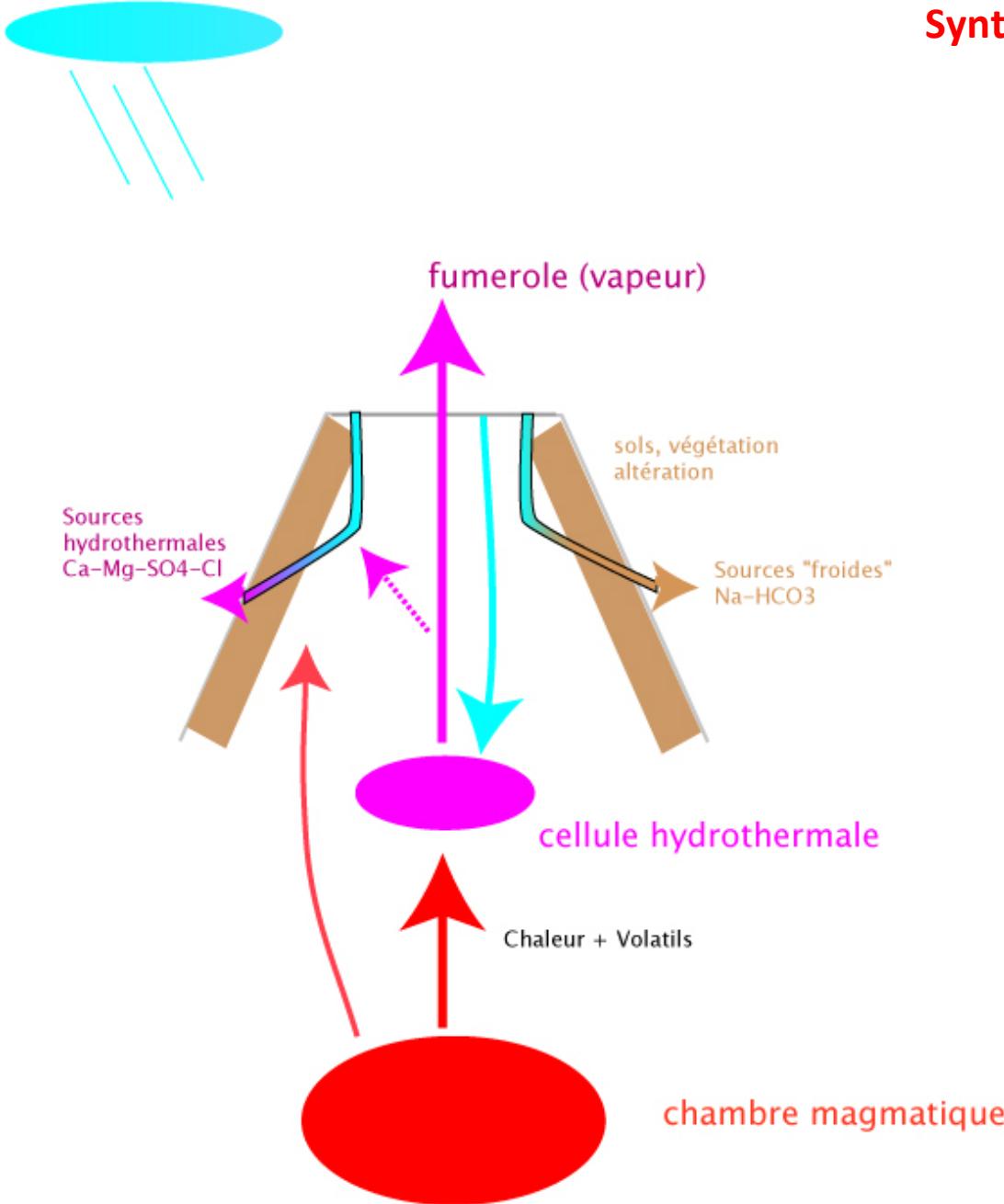
Rapports isotopiques de l'Hélium des sources

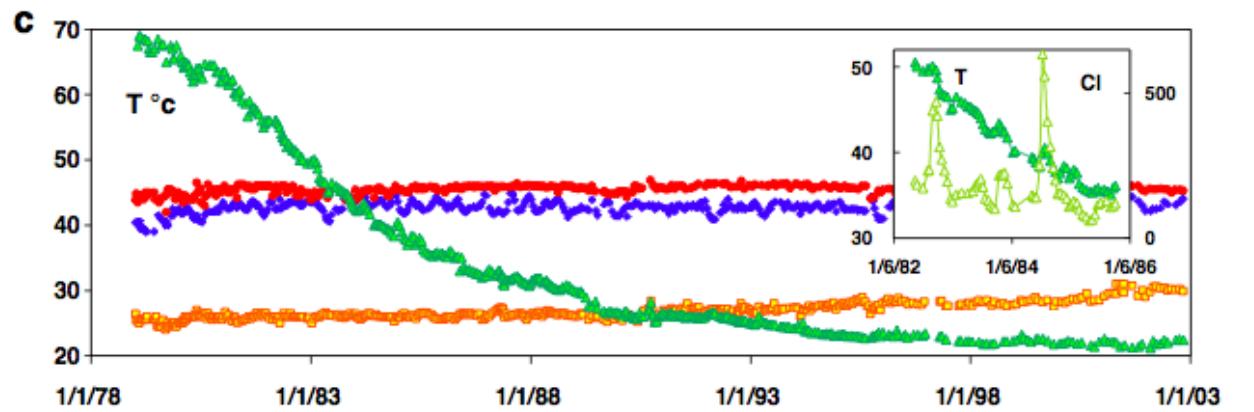
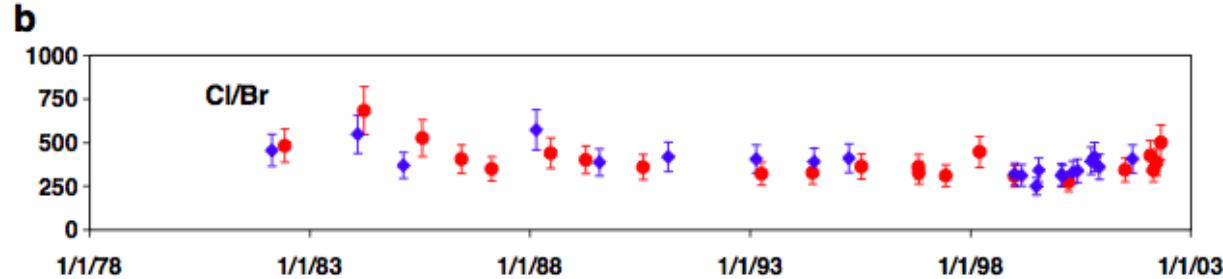
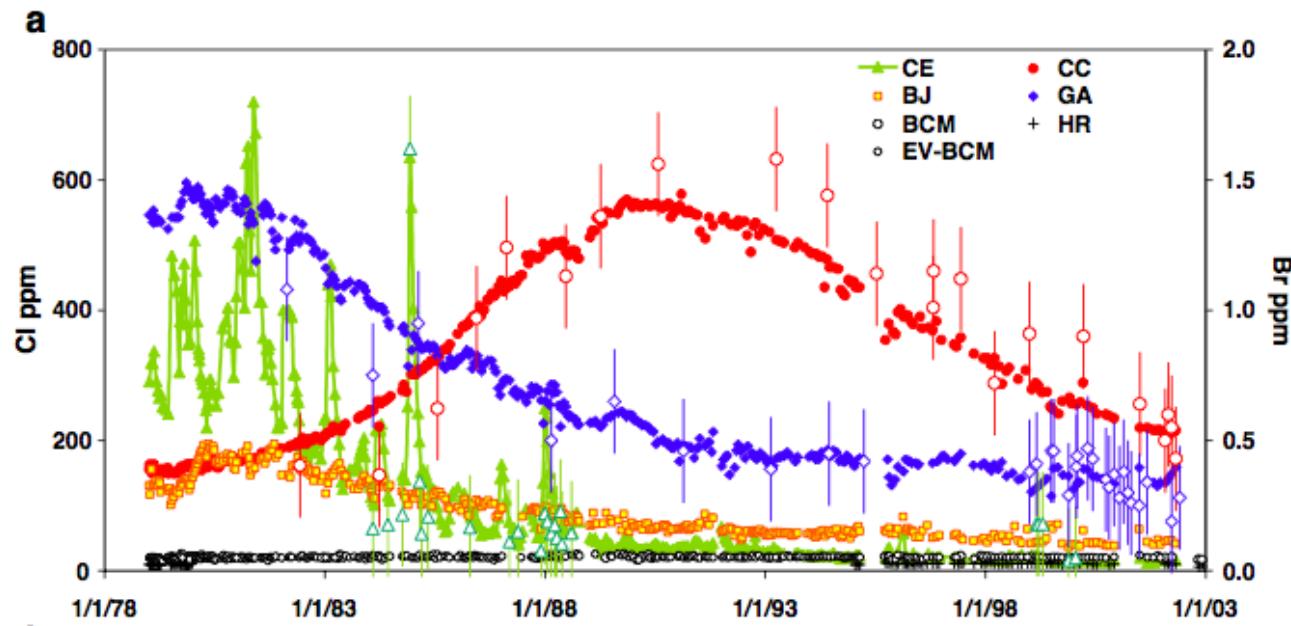




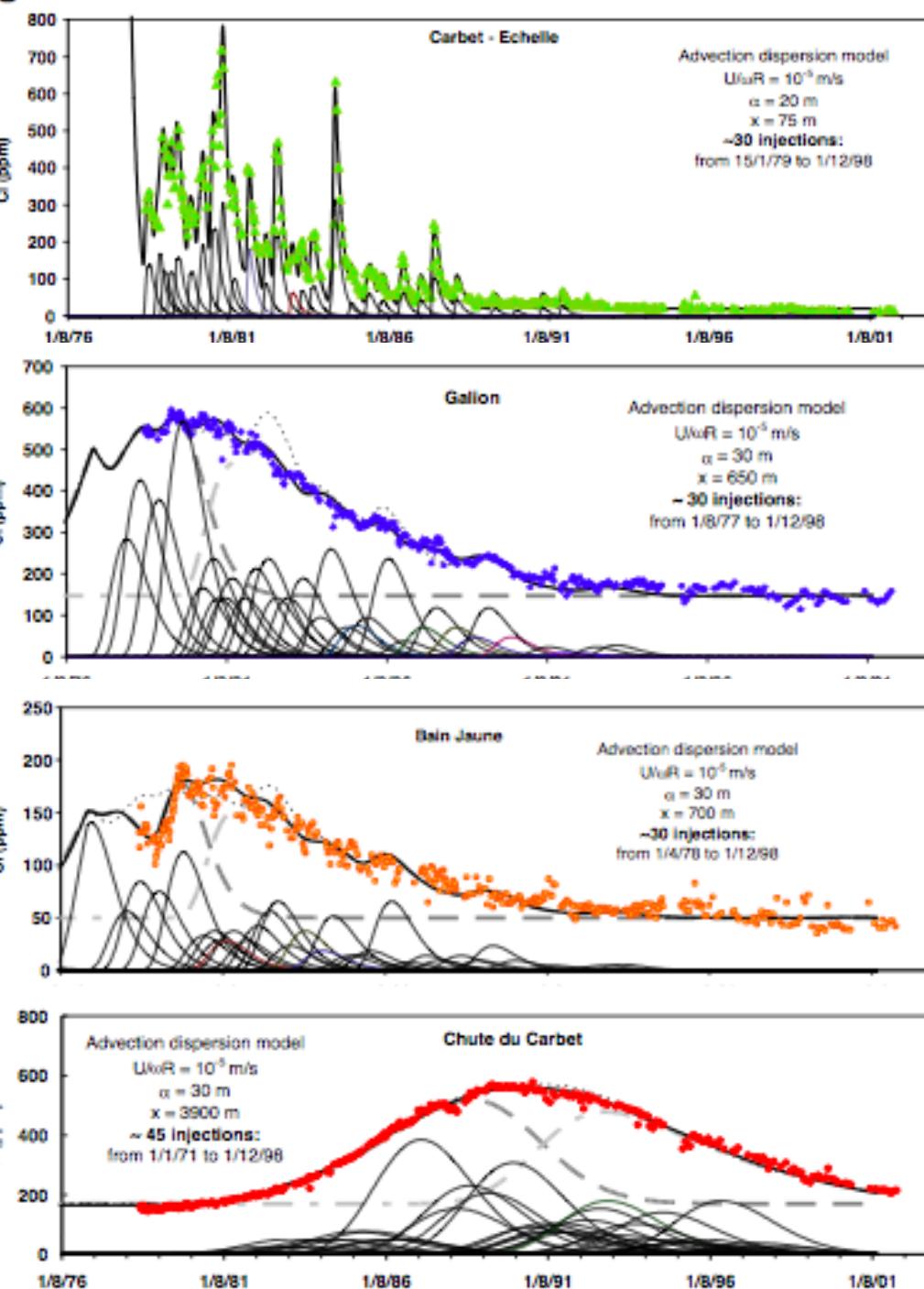
Pluie

Synthèse Hydrogéochimique





Villemant et al, 2005

C

Propagation des
Anomalies de chlorures

Distance augmente

Villemand et al., 2005

Isotopes du Chlore et du brome

$\text{Cl}^{35} \approx 75.7\%$, $\text{Cl}^{37} \approx 24.3\%$

$\text{Br}^{79} \approx 50.7\%$, $\text{Br}^{81} \approx 49.3\%$

Notation

$$R = \text{Cl}^{37}/\text{Cl}^{35} \approx 0.321 \quad \text{SMOC}$$

$$\delta^{37}\text{Cl} = (R_{\text{sample}}/\text{SMOC} - 1) * 1000 \quad \text{\textperthousand}$$

$$R = \text{Br}^{81}/\text{Cl}^{79} \approx 0.97 \quad \text{SMOB}$$

$$\delta^{81}\text{Br} = (R_{\text{sample}}/\text{SMOB} - 1) * 1000 \quad \text{\textperthousand}$$

$$\text{facteur de fractionnement} = \alpha_{A-B} = R_A/R_B$$

Paysage isotopique

Chlore: -8 ‰ (fluides de la CO) to 10 ‰ (HCl des gaz volcaniques)

Brome : ? Quelques ‰

pourquoi ?

Cl⁻ and Br⁻ : non reactifs

⇒ $\delta^{37}\text{Cl}$ et $\delta^{81}\text{Br}$: traîneurs isotopiques des sources mélanges et des processus de transport (diffusion, advection, ...).

Séquestration géologique du CO₂

$\delta^{37}\text{Cl}$ et $\delta^{81}\text{Br}$ pourraient détecter des modifications de

- permeabilité/porosité dans les aquifères
- Barrière de perméabilité

Protocole analytique pour les isotopes du chlore

Goal : 20 micromoles of CH₃Cl : Kaufman, 1984; Eggenkamp, 1994

1. add AgNO₃ and pH buffer => AgCl precipitate
2. react AgCl with CH₃I in excess to get CH₃Cl
3. purify CH₃Cl by gas chromatography
4. Dual inlet or He continuous flow mass spectrometry

ion mass 50 : CH₃Cl³⁵

ion mass 52 : CH₃Cl³⁷



$\delta^{37}\text{Cl}$

Pour les isotopes du brome

Goal : 20 micromoles of CH_3Br : Eggenkamp, 2003

0. get Br/Cl ratio to ≈ 1 by oxidative distillation
1. add AgNO_3 and pH buffer => AgBr precipitate
2. react AgBr with CH_3I in excess to get CH_3Br
3. purify CH_3Br by gas chromatography
4. Dual inlet or He continuous flow mass spectrometry

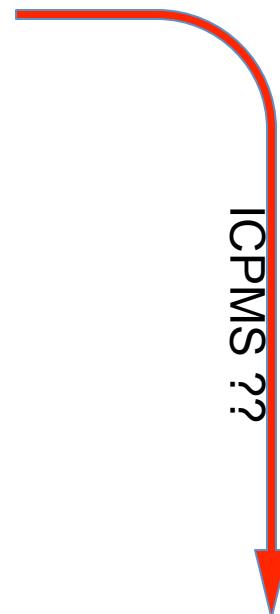
ion mass 94 : $\text{CH}_3\text{Br}^{79}$

ion mass 96 : $\text{CH}_3\text{Br}^{81}$

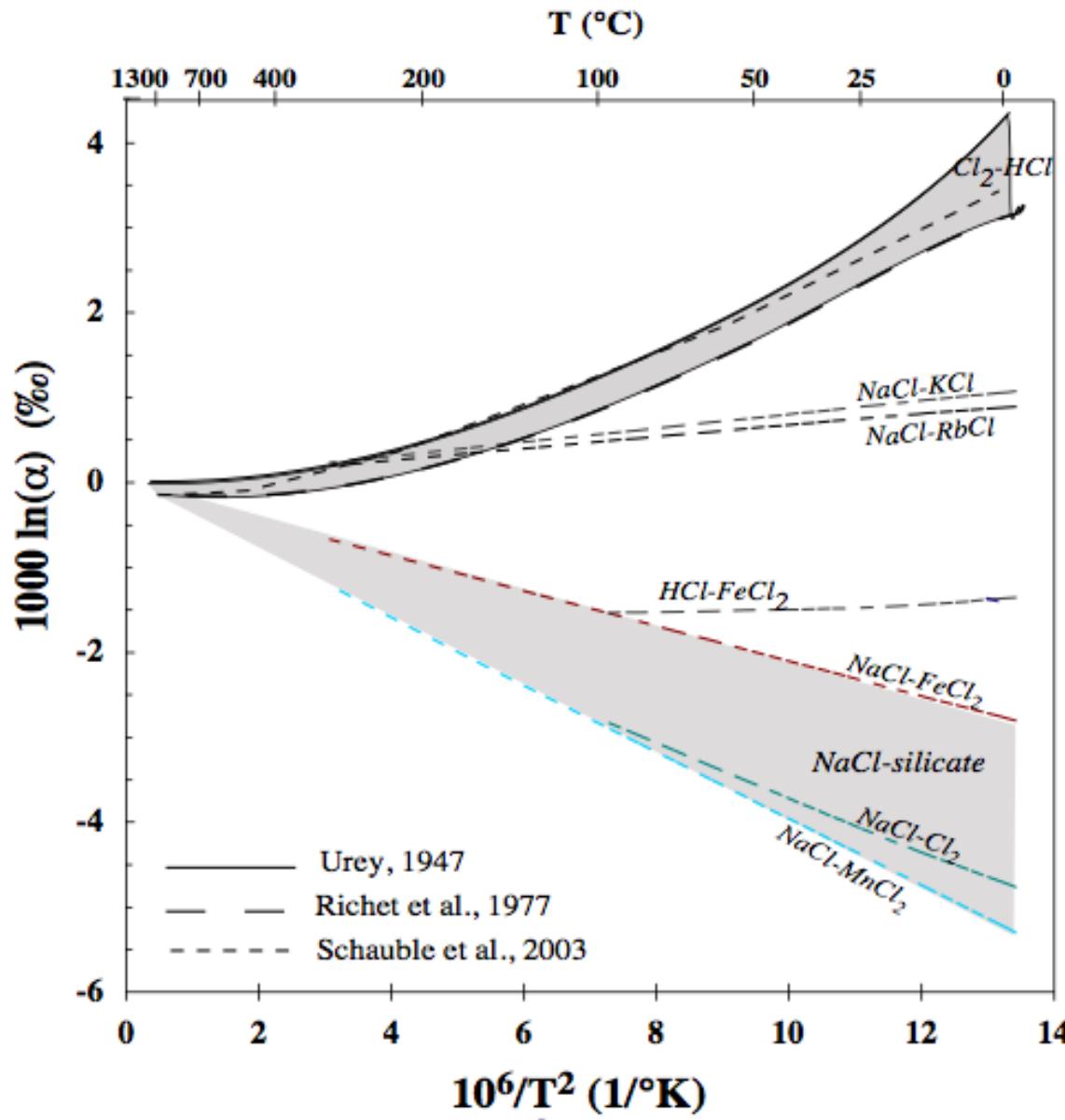


$\delta^{81}\text{Br}$

ICPMS ??

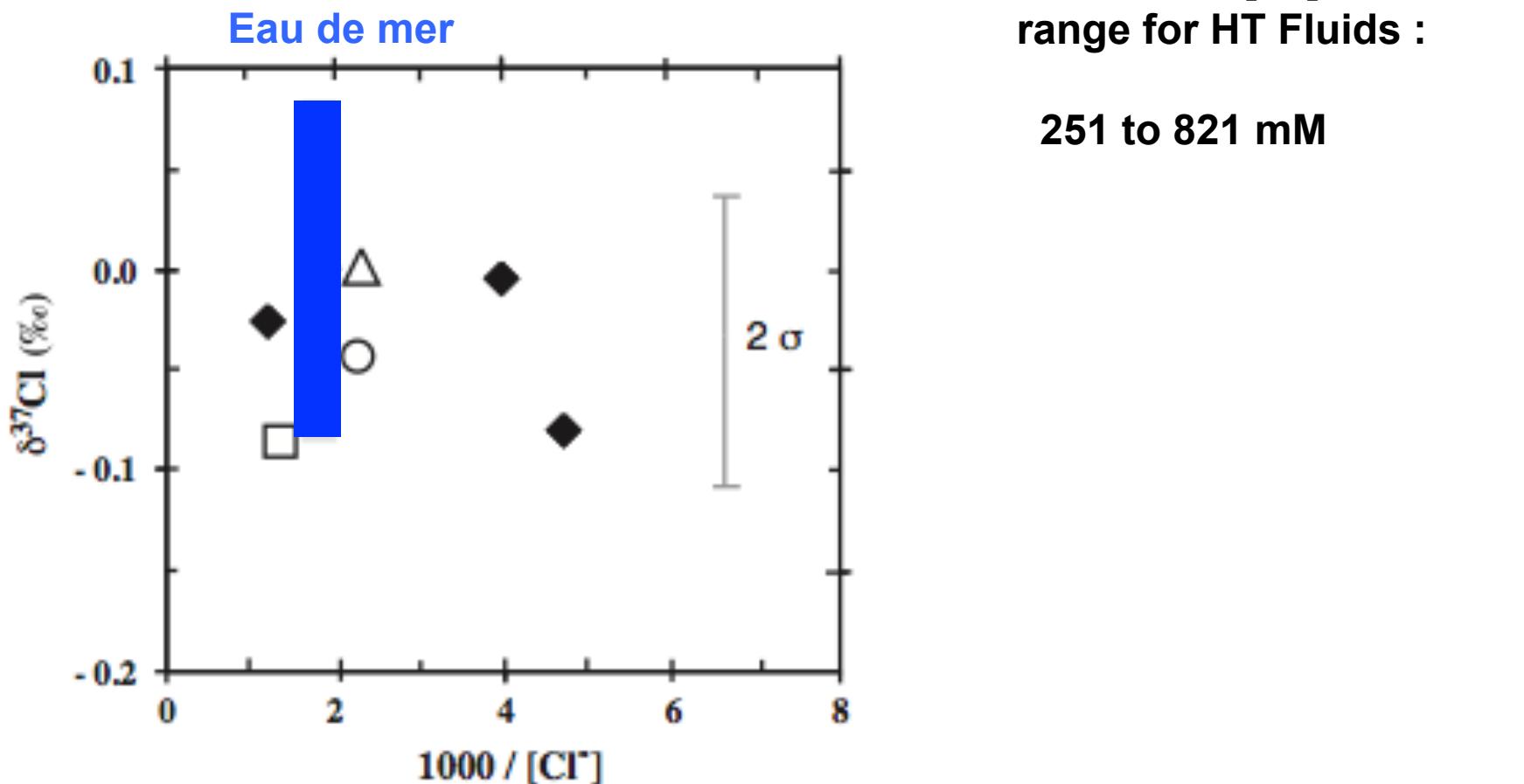


Fractionnements isotopiques à l'équilibre thermodynamique



Fluides hydrothermaux des rides océaniques

T > 200°C

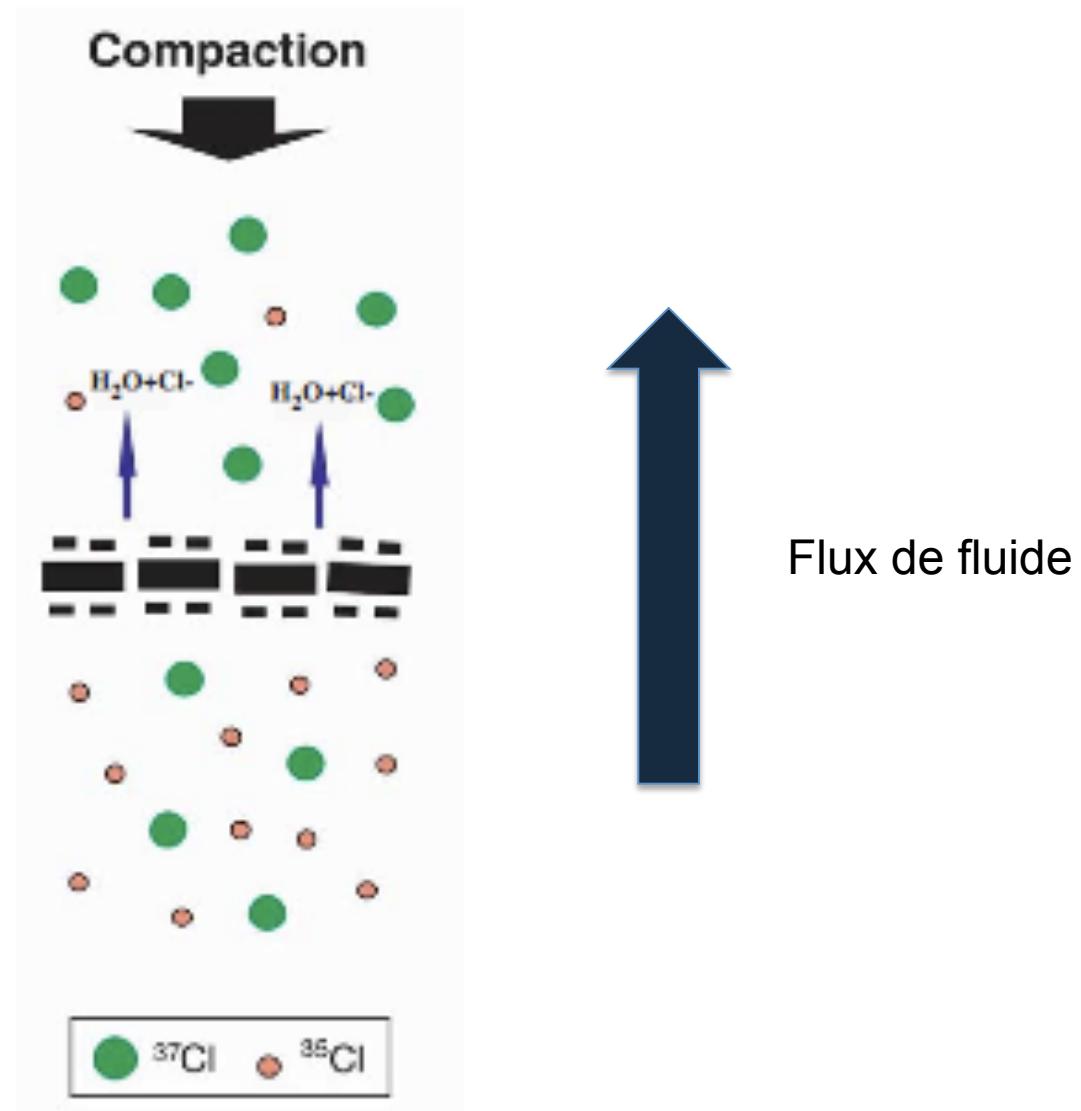


Théorie : la filtration ionique de Phillips et Bentley 1987

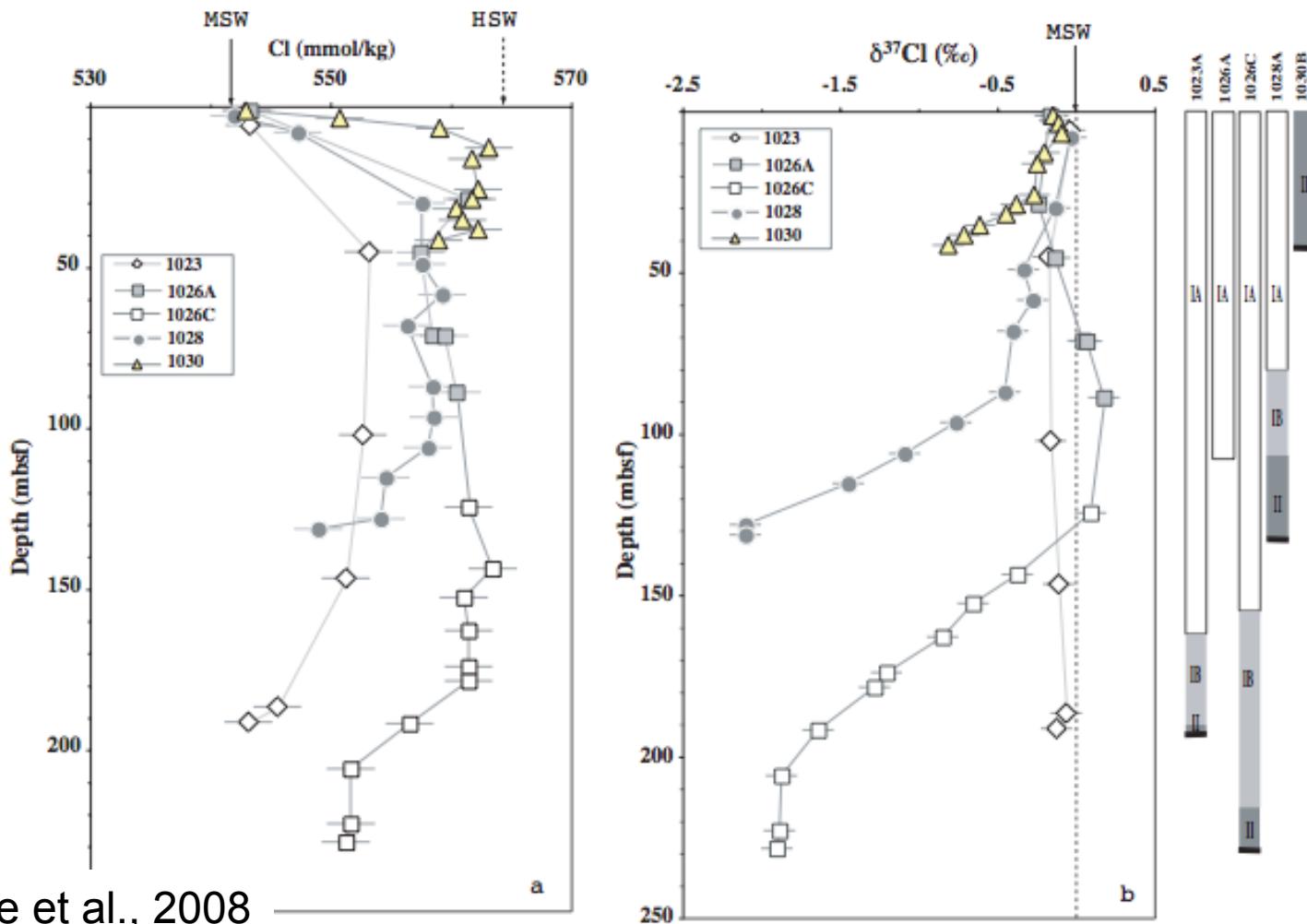
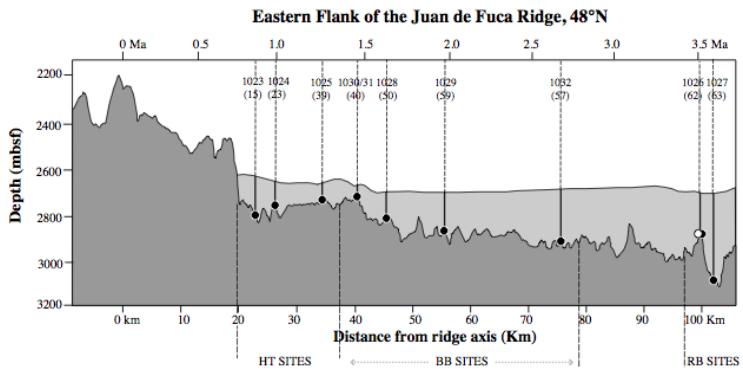
$\delta^{37}\text{Cl}$ augmente
[Cl] décroît

Membrane
Semi permeable
Chargée négativement

$\delta^{37}\text{Cl}$ diminue
[Cl] augmente

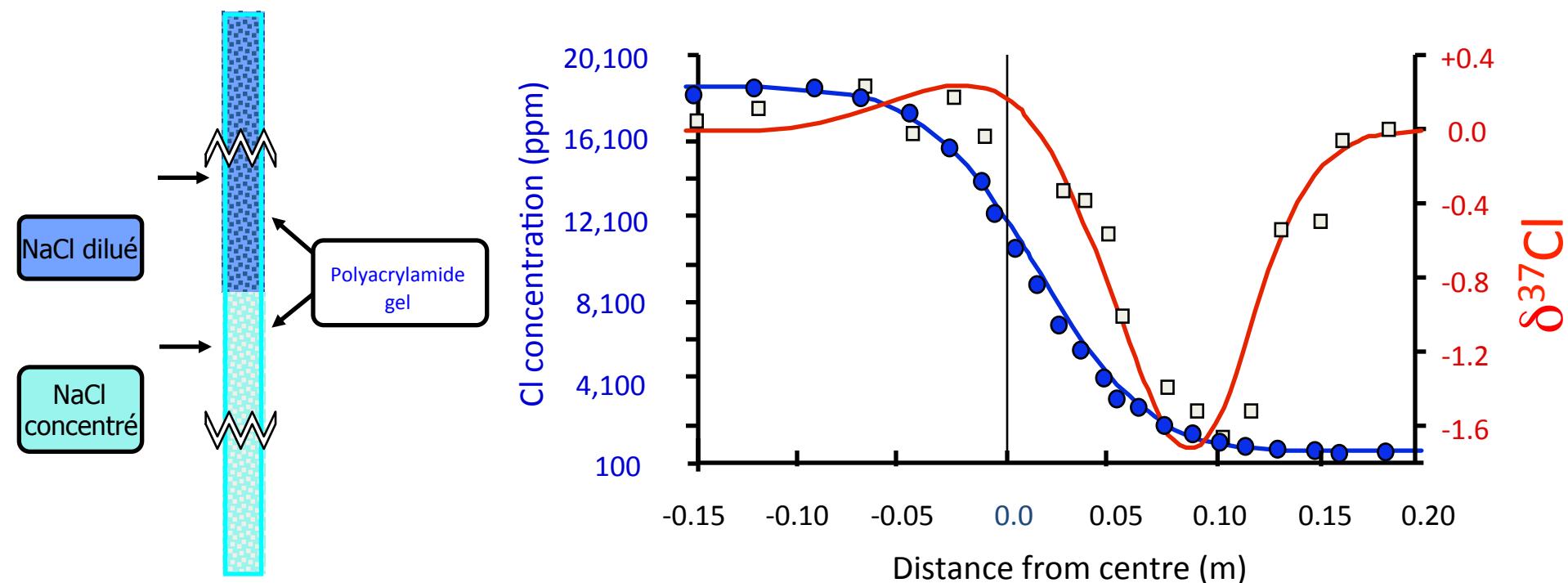


Fluides des sédiments Du flanc de la dorsale Juan de Fuca Ridge

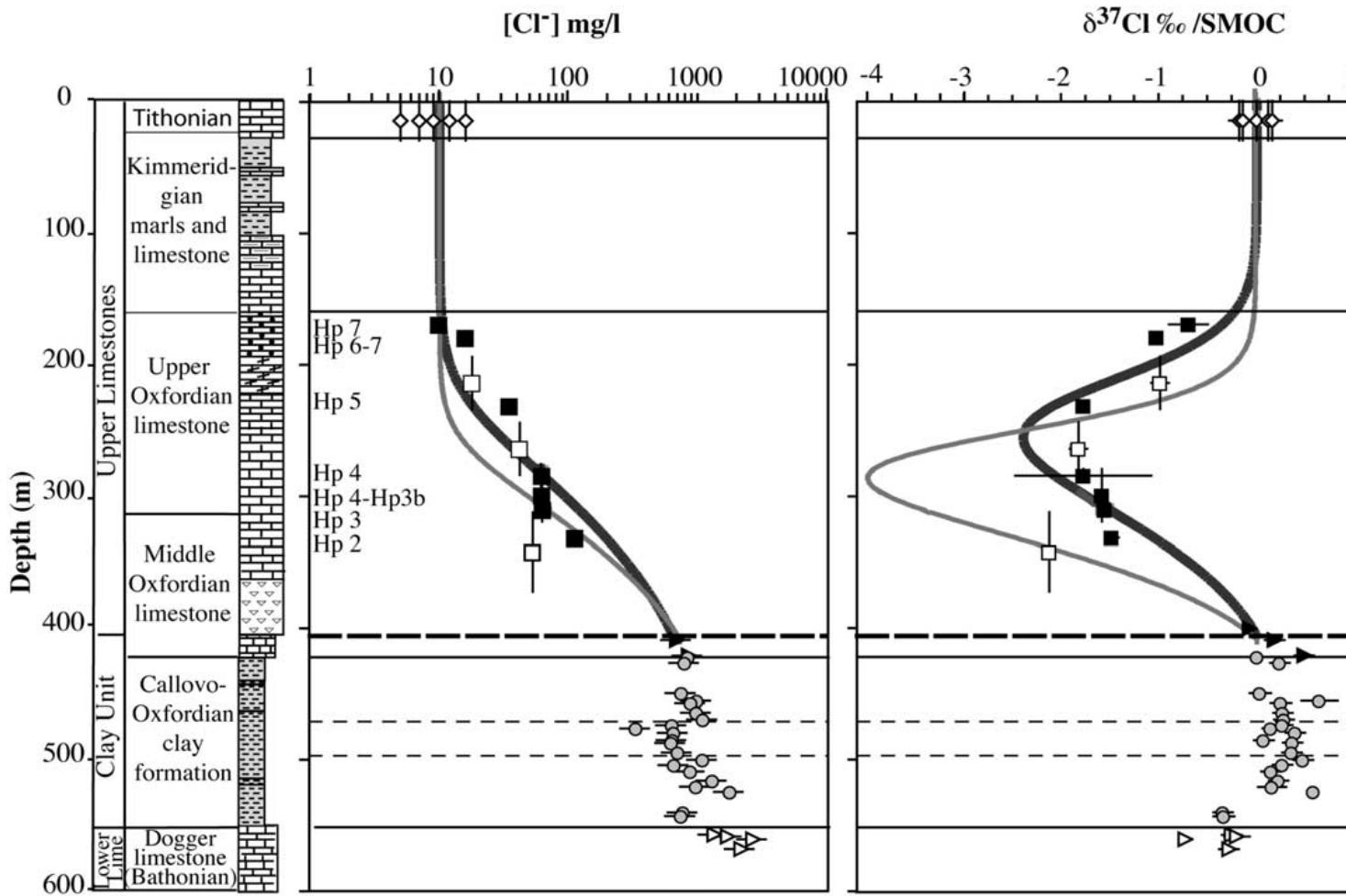


Fractionnement des isotopes du chlore par la diffusion

- Diffusion de Cl^- et Br^- dans un gel de Polyacrylamide
 - $D_{35\text{Cl}}/D_{37\text{Cl}}$ (21°C) : 1.00166
 - $D_{79\text{Br}}/D_{81\text{Br}}$ (21°C) : 1.00064



Fuite des chlorures dans le bassin de Paris



Upper Limestones water samples : ◊ Tithonian ■ Oxfordian-underground laboratory area □ Oxfordian- EST 311 site

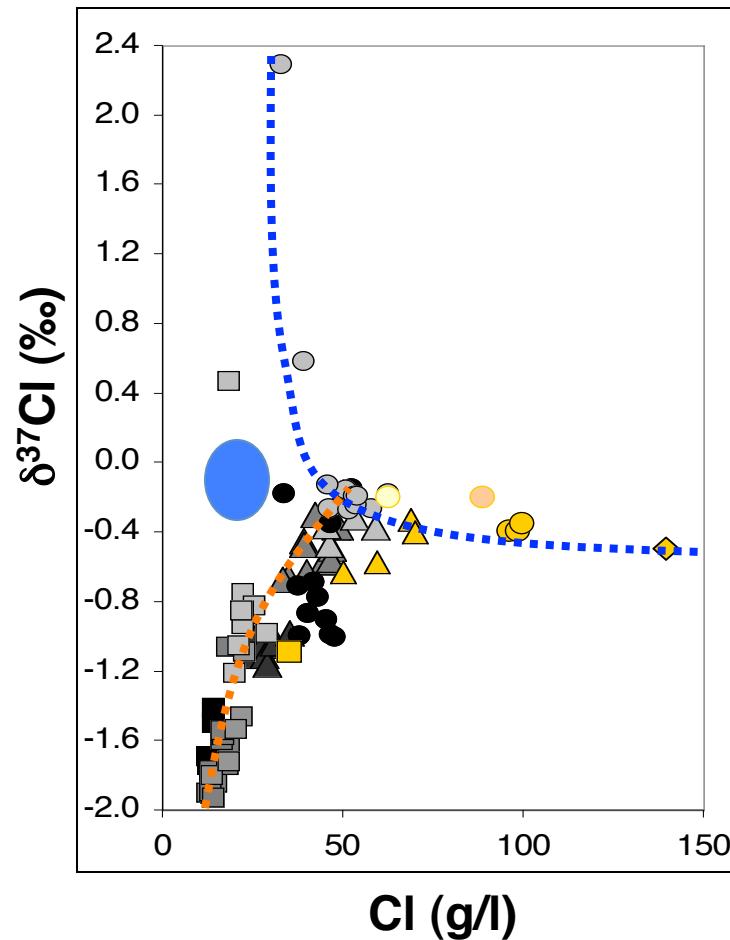
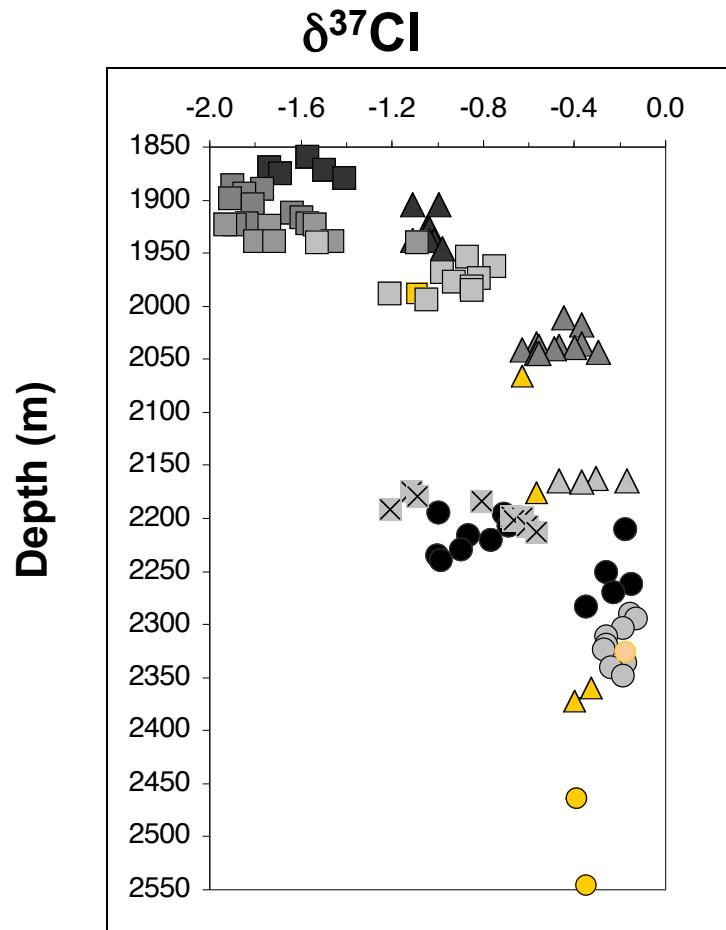
Clay Unit and Lower Limestone rock samples (HTM 102) : ▶ basal 20m Middle Oxfordian ○ Callovo-Oxfordian ▶ I

— lithological limit ----- limit of the central part of the Callovo-Oxfordian clay —— diffusion boundary

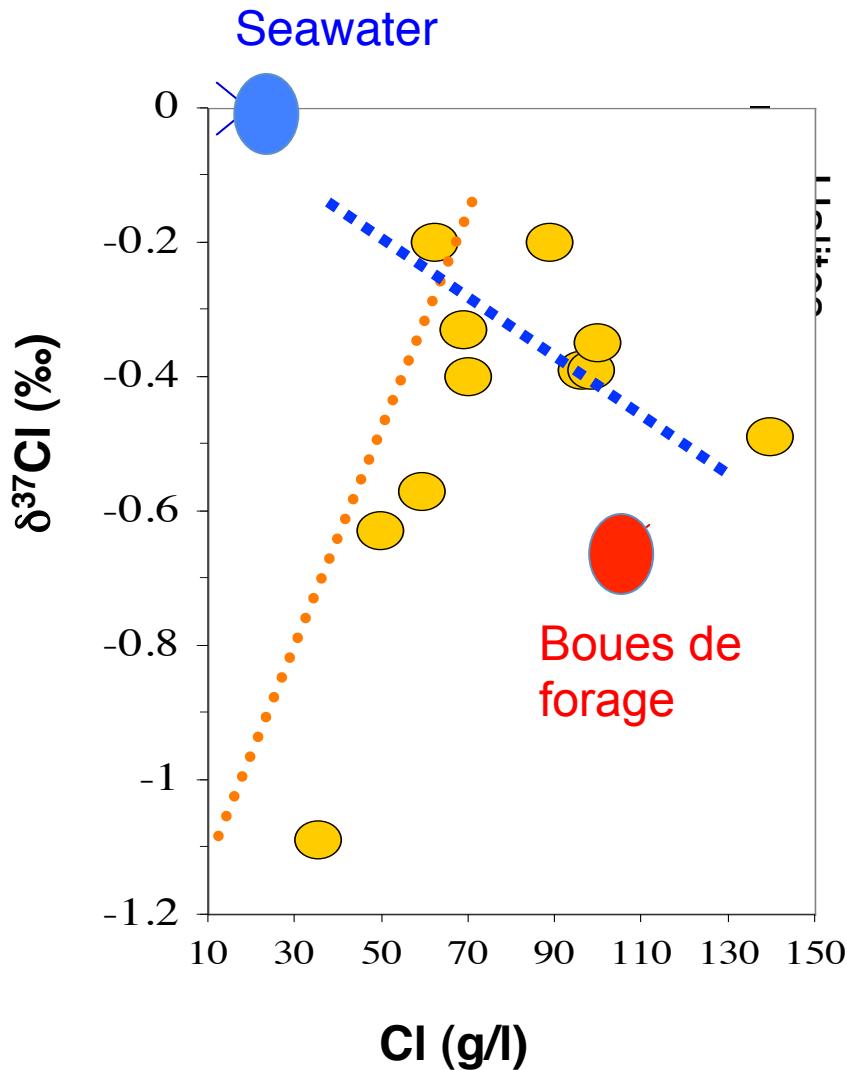
■ Chlorine transport model for URL site (◊, ■) — Chlorine transport model for EST311 site (◊, □)

Lavastre et al., 2005

$\delta^{37}\text{Cl}$ des saumures des réservoirs pétroliers



Interpretation



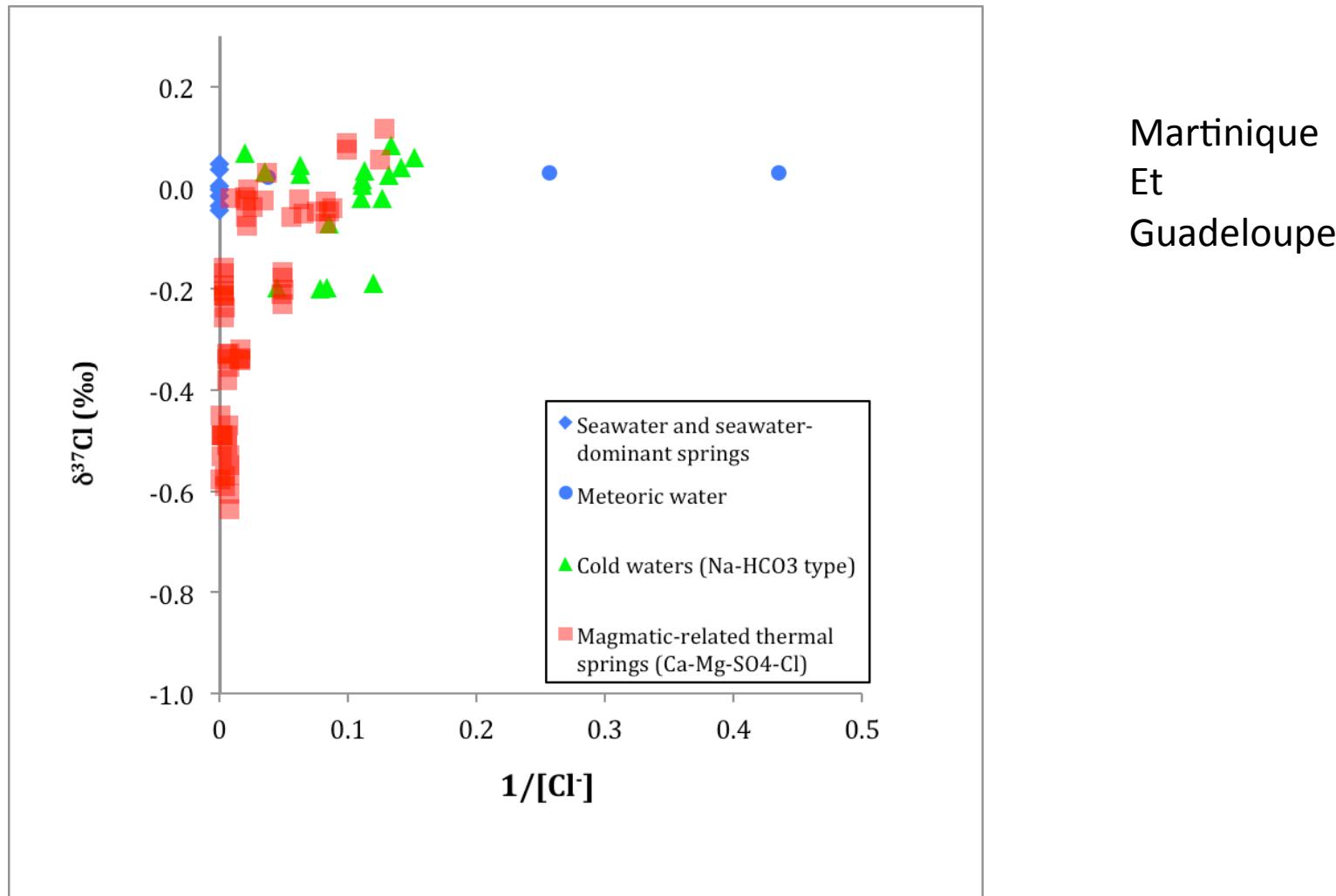
Filtration ionique
mélange

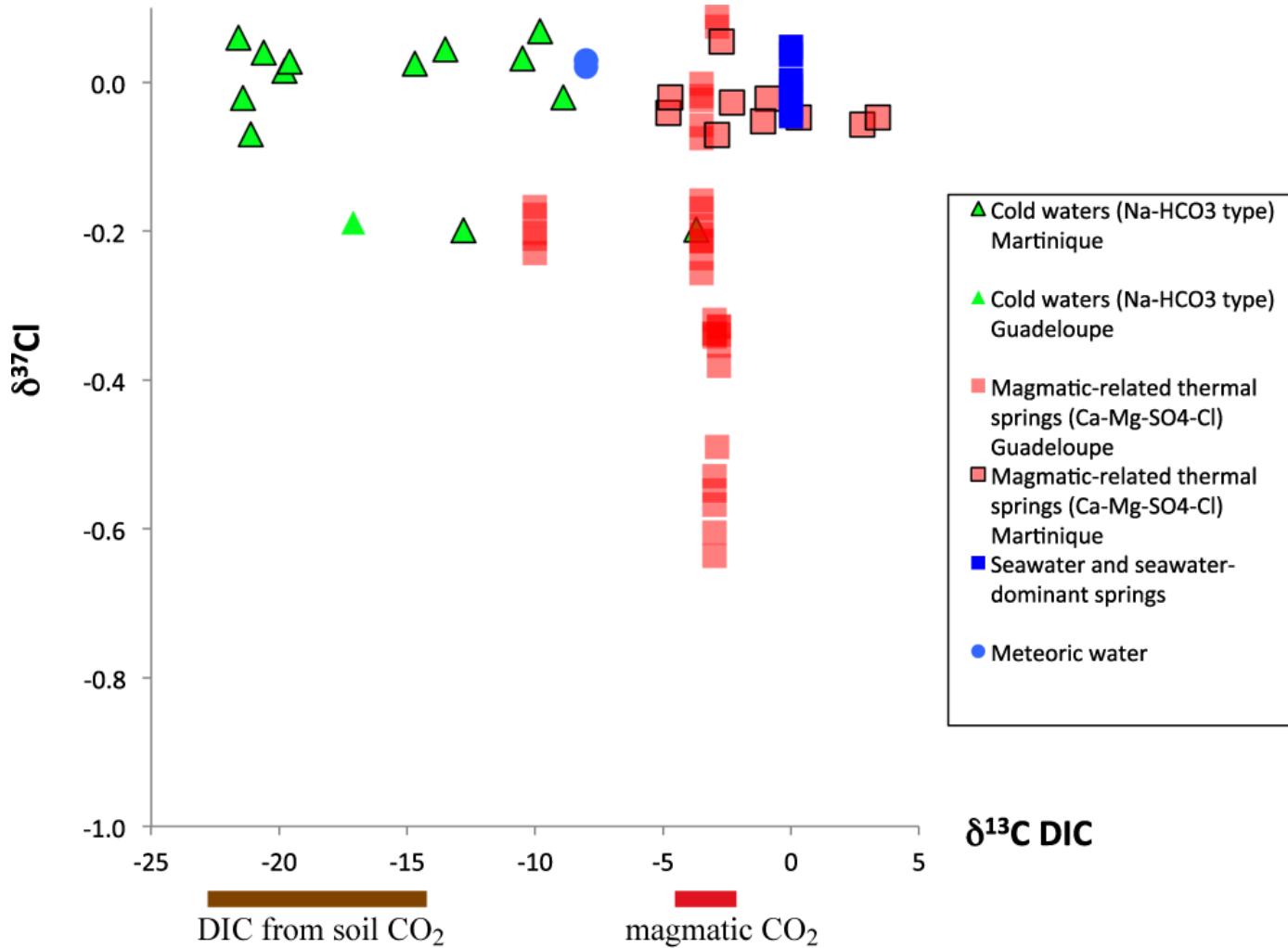
+

Pas de contribution
des evaporites,
Ni des boues de forages

Woulé Ebongué

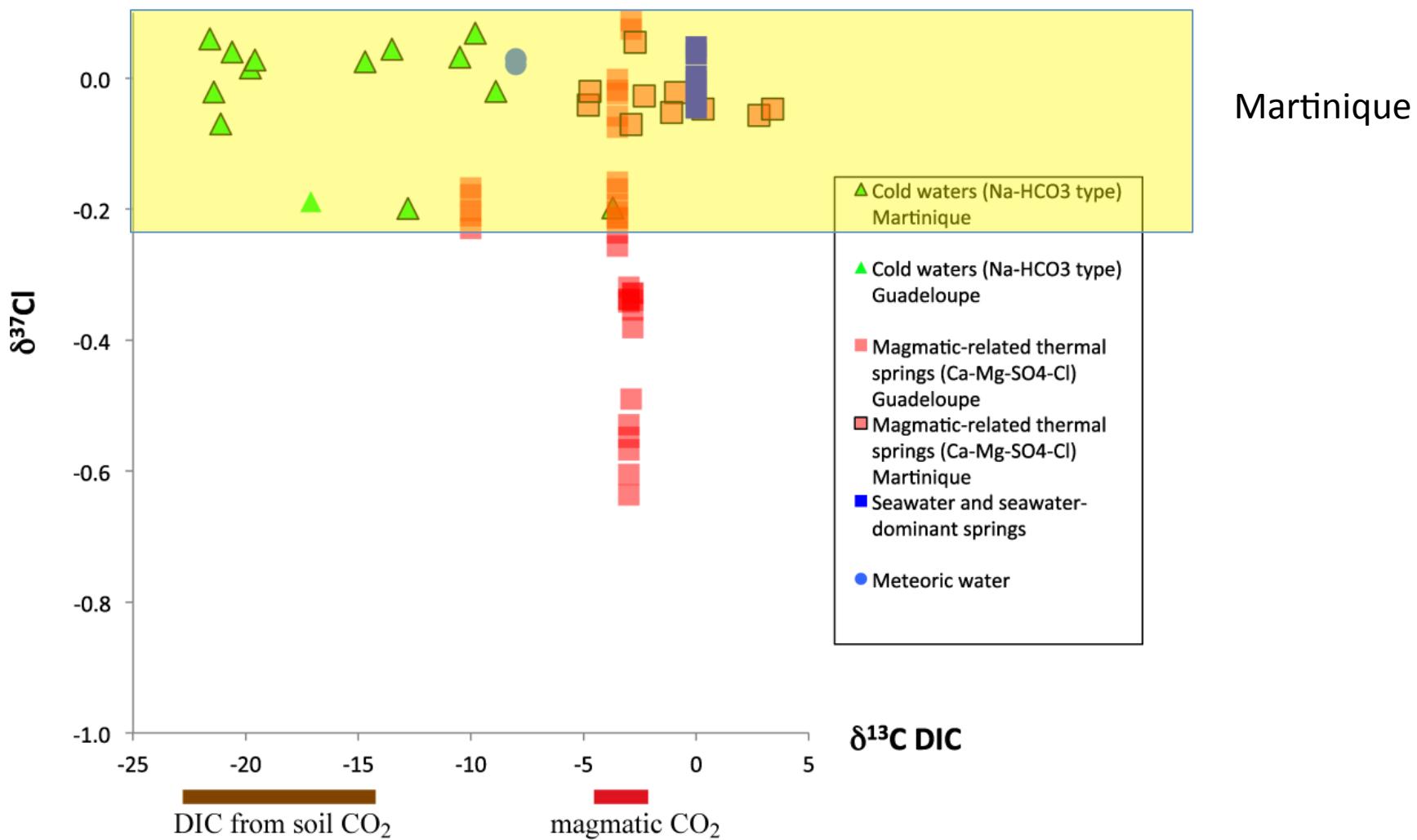
D'où viennent les chlorures des sources hydrothermales ?





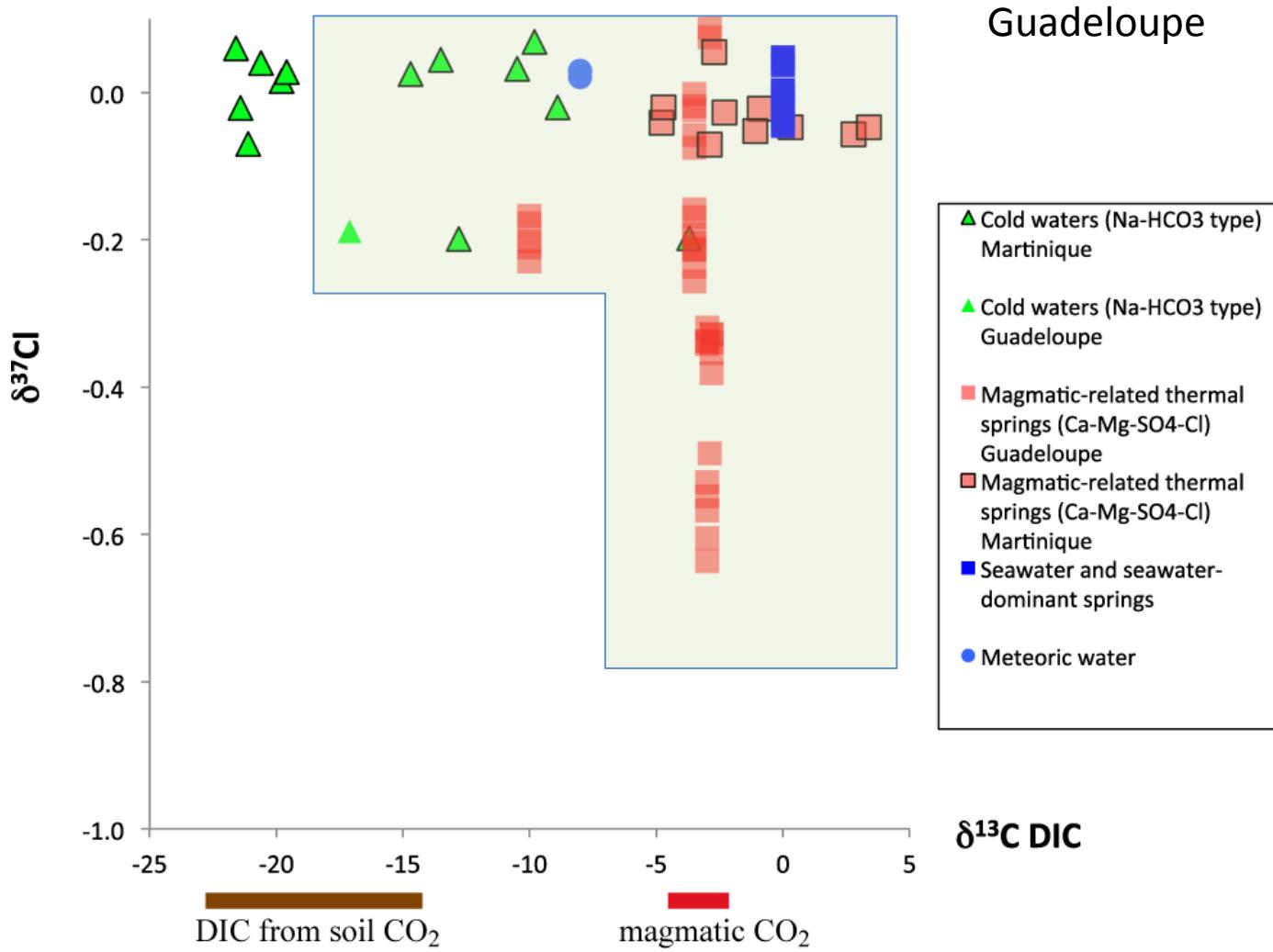
Chlore magmatique : $\delta^{37}\text{Cl} < -0.2 \text{ ‰}$

Chlore de « surface » : $\delta^{37}\text{Cl} \geq -0.2 \text{ ‰}$



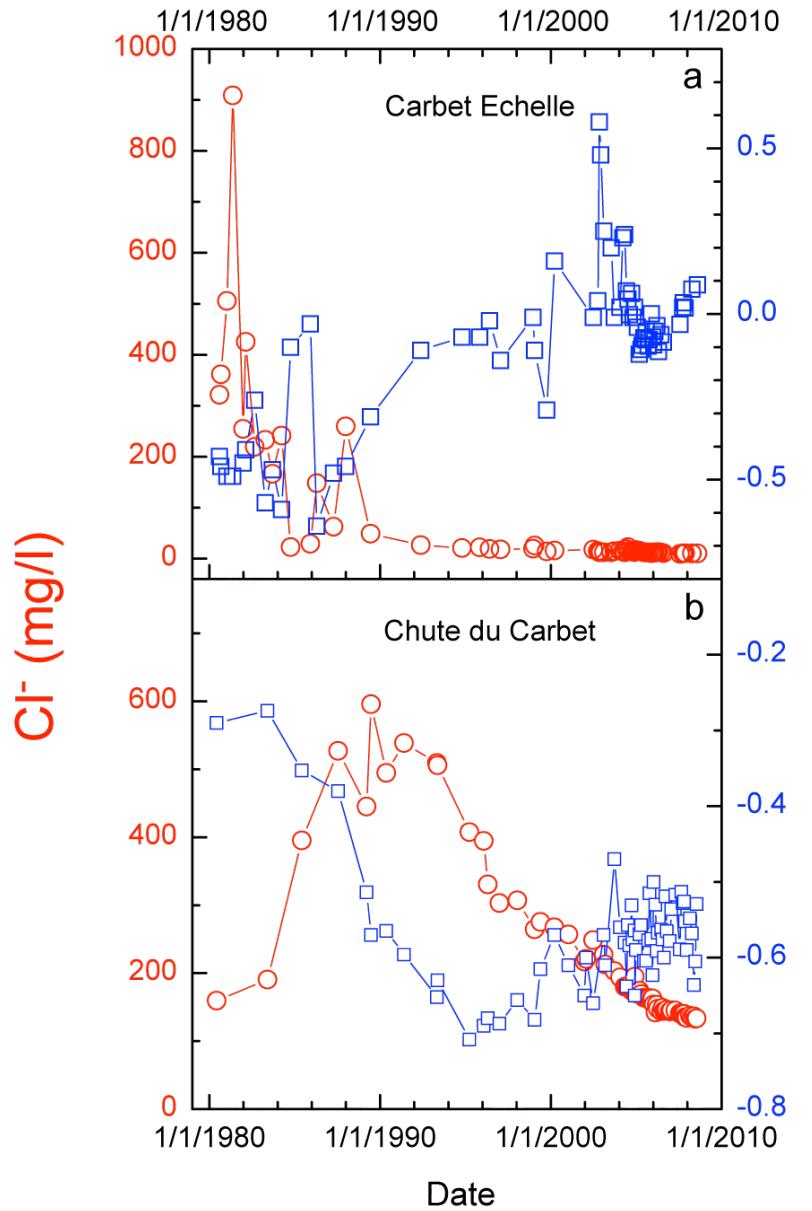
Chlore magmatique : $\delta^{37}\text{Cl} < -0.2 \text{ ‰}$

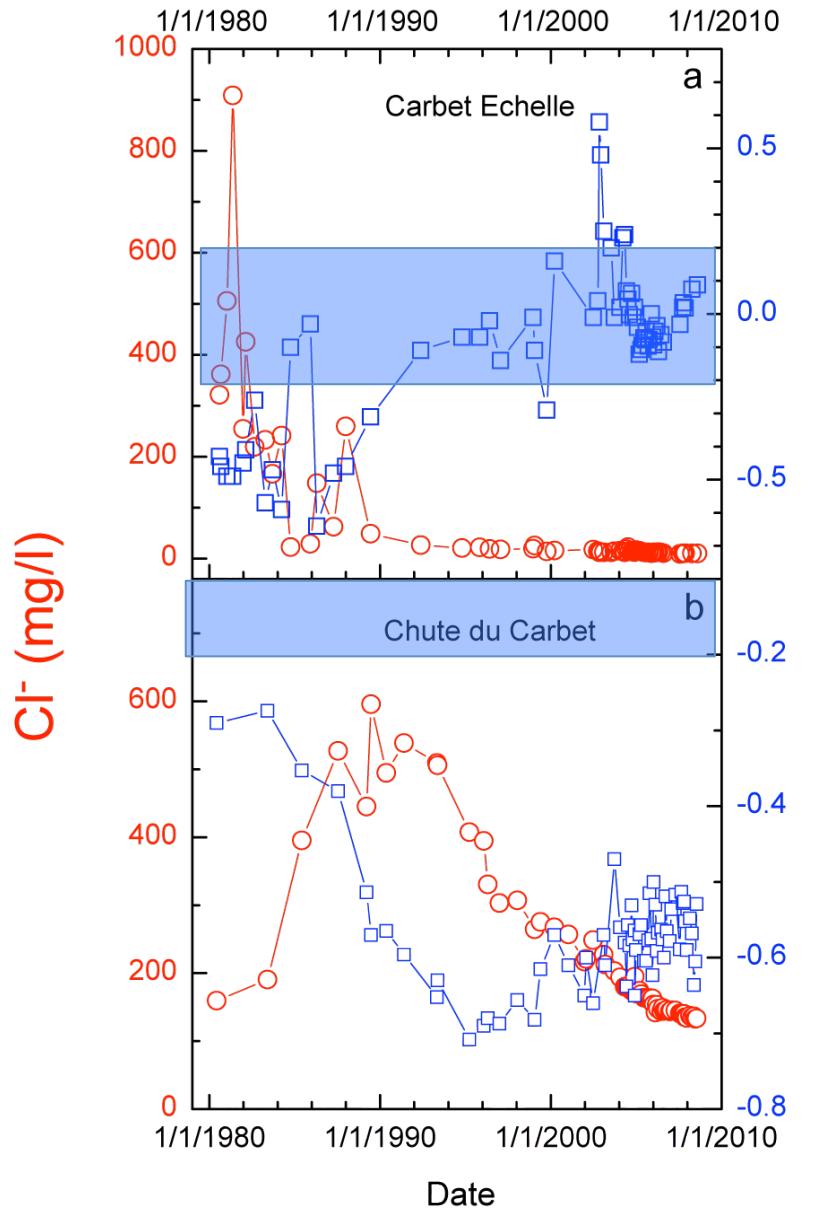
Chlore de « surface » : $\delta^{37}\text{Cl} \geq -0.2 \text{ ‰}$



Chlore magmatique : $\delta^{37}\text{Cl} < -0.2 \text{ ‰}$

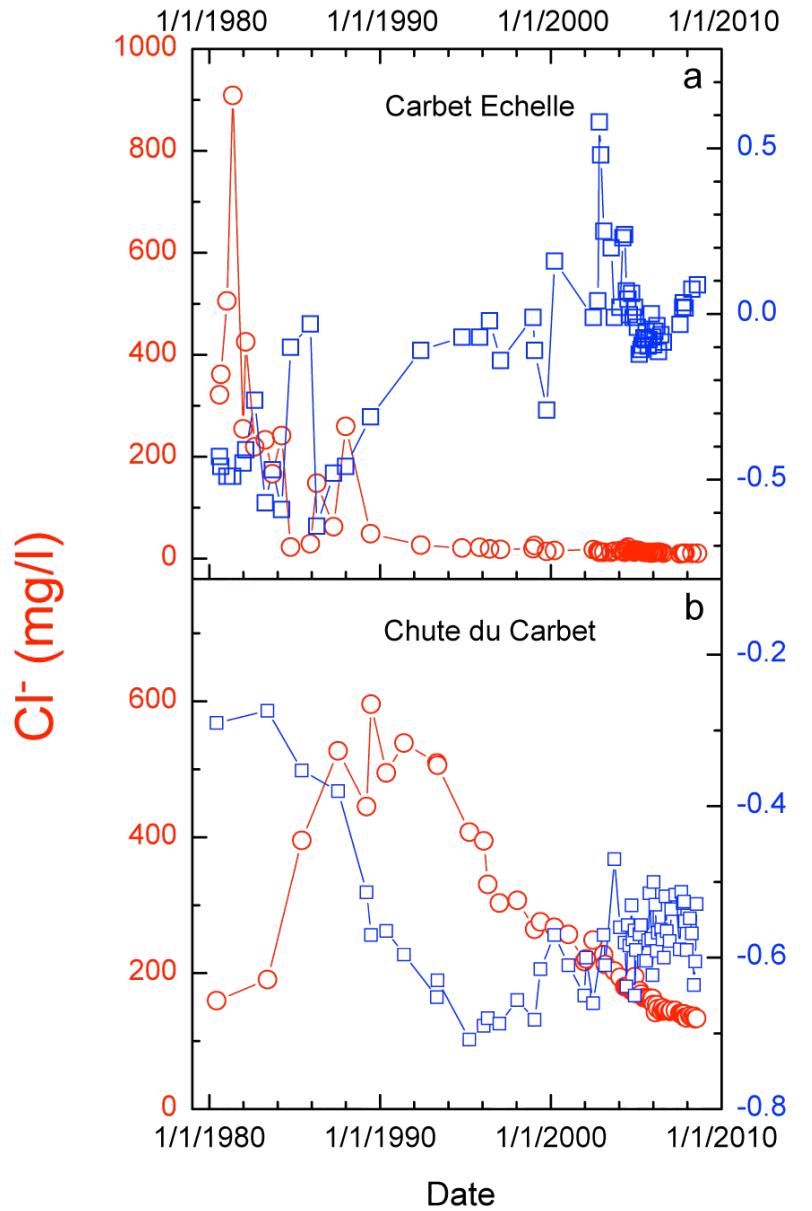
Chlore de « surface » : $\delta^{37}\text{Cl} \geq -0.2 \text{ ‰}$





CE : plus de Cl magmatique

CC : toujours du chlore magmatique



- Montagne Pelée : $\delta^{37}\text{Cl} \geq -0.2\text{ ‰}$
- Soufrière : $\delta^{37}\text{Cl} < -0.2\text{ ‰}$

Temps de résidence du chlore court

- 80 ans plus de Cl magmatique
- 35 ans tjs du Cl magmatique

